

# *DREAM: Dynamical Research Empirical Atmospheric Model*

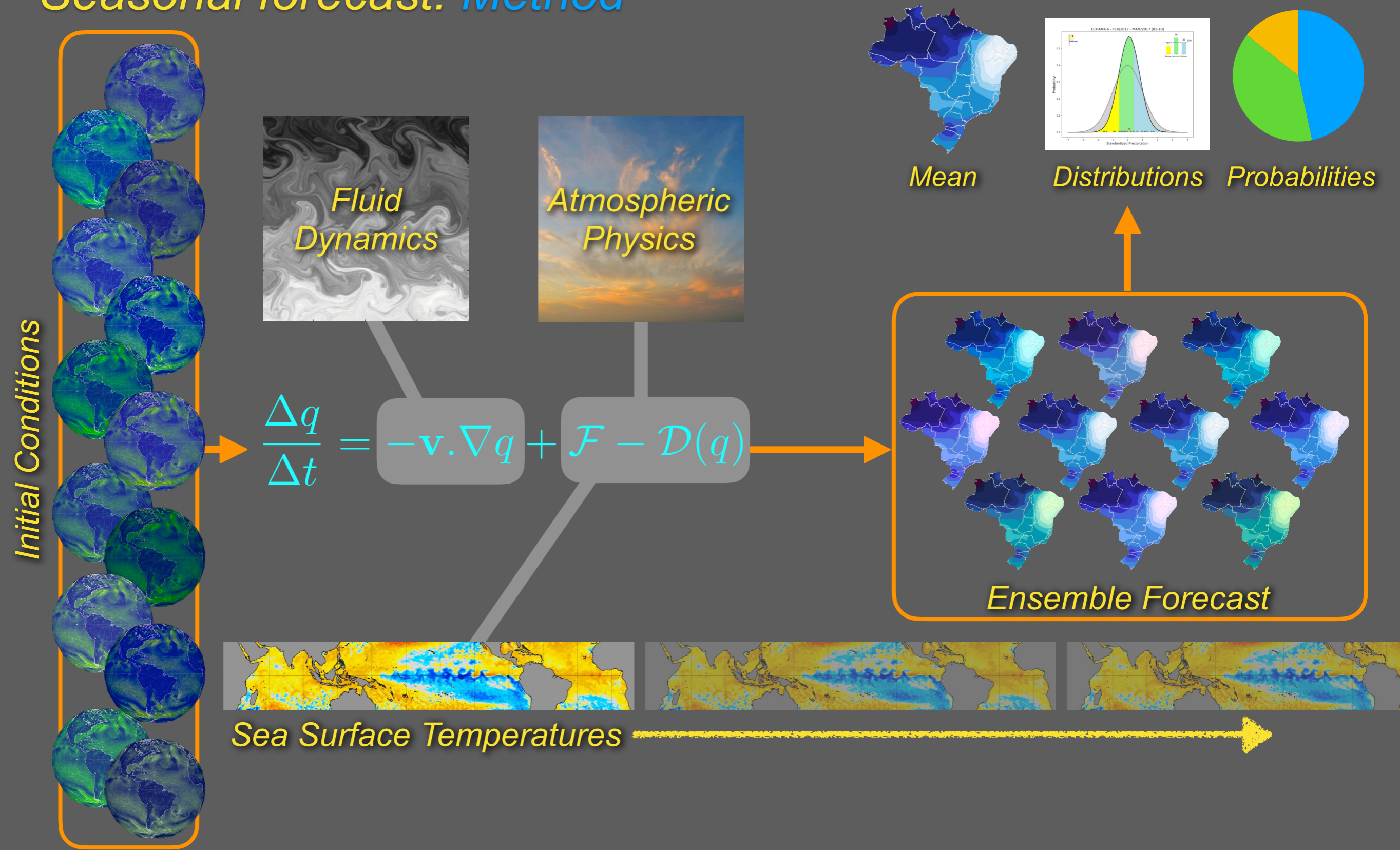
*Two questions:*

*1) How to we replace the explicit physics in a global general circulation model with data-derived alternative ?*

*2) How do we represent the relationship between sea surface temperatures and precipitation over the ocean ?*



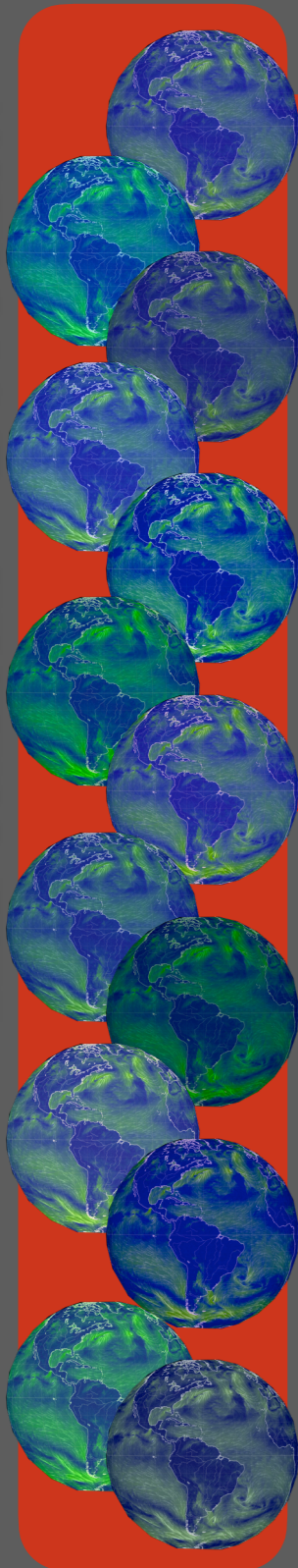
# Seasonal forecast: Method



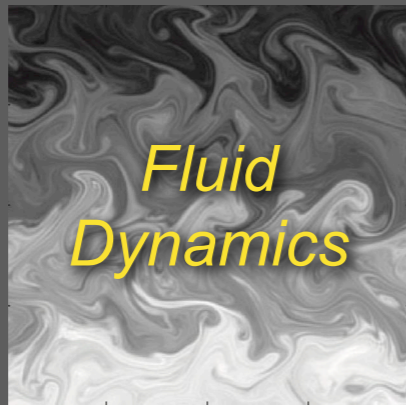
ECHAM 4.6 Global GCM, Resolution T42L19. 20-member ensemble. RSM97 Regional Downscaling model.  
One-month single member forecast: ~ **ONE HOUR** CPU (FUNCEME computer system)

# Seasonal forecast: *DREAM*

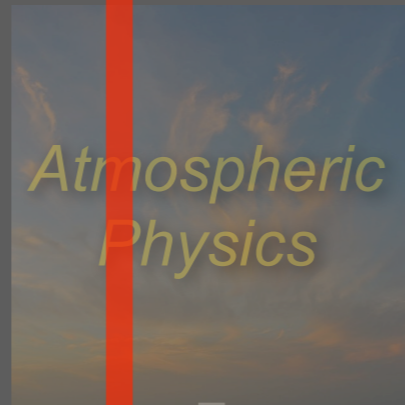
Initial Conditions  
ERA-Interim Dataset



Empirical Forcing

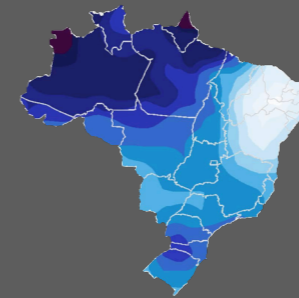


Fluid Dynamics

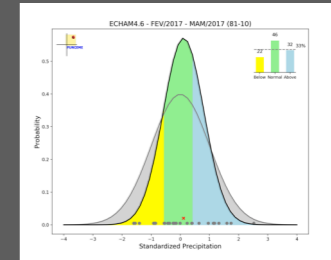


Atmospheric Physics

$$\frac{\Delta q}{\Delta t} = -\mathbf{v} \cdot \nabla q + \mathcal{F} - \mathcal{D}(q)$$



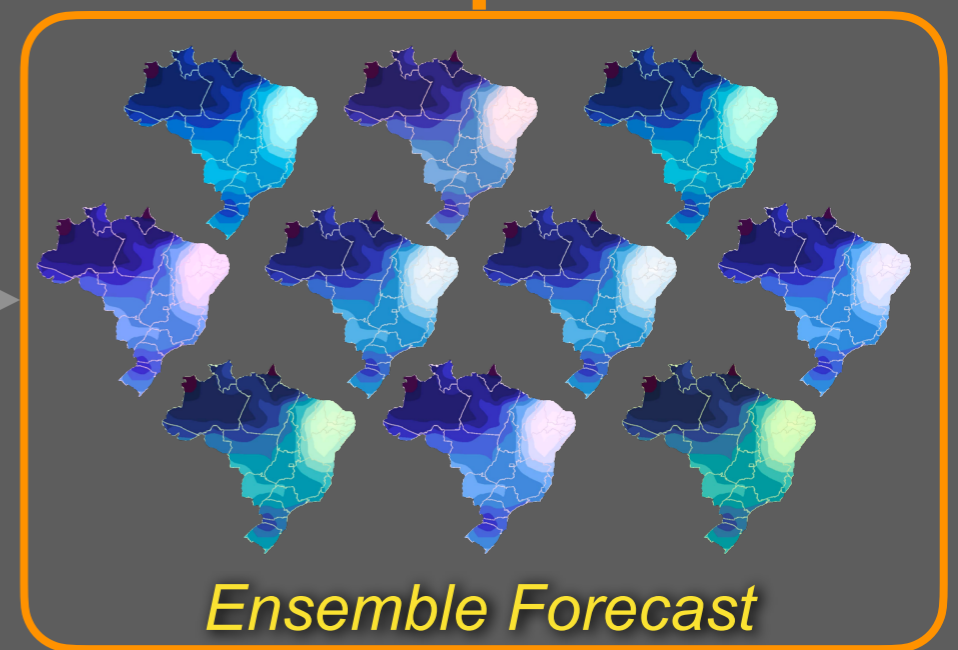
Mean



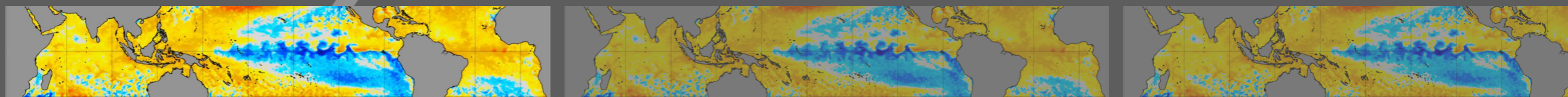
Distributions



Probabilities



Ensemble Forecast



Sea Surface Temperatures

*DREAM* Global Dynamical Model, Resolution T42L15. 38-member ensemble.  
One-month single member forecast: ~ **ONE MINUTE** CPU (this laptop).

# Empirical forcing philosophy

Consider the development of potential vorticity,  $q$  in the atmosphere:

$$\frac{\partial q}{\partial t} + \mathbf{v} \cdot \nabla q = \mathcal{F}(t) - \mathcal{D}(q)$$

*External Forcing*

Make a model of this by assuming the external forcing is fixed

Let's construct a model of this.  
The only thing external to our model is  $F$   
Substitute a forcing  $G$  deduced from data.

$$\frac{\partial q}{\partial t} + \mathbf{v} \cdot \nabla q + Dq = \mathcal{G}$$

We can use the model to deduce the forcing  $G$  from a time series of reanalysis data.

We can calculate  $G$  by taking the average of a large number of single-timestep integrations of the unforced model

$$\begin{aligned} \mathcal{G} &= \overline{\mathcal{F}} = \overline{\mathbf{v} \cdot \nabla q} + D\bar{q} \\ &= \frac{1}{n} \sum_{i=1}^n \mathbf{v}_i \cdot \nabla q_i + Dq_i = -\frac{1}{n} \sum_{i=1}^n \left. \frac{\partial q_i}{\partial t} \right|_{uf} \end{aligned}$$

Add the intelligence:

$$\mathcal{F}(t) \rightarrow \mathcal{F}(t, q) \quad \mathcal{G} \rightarrow \mathcal{G}(q)$$

# Condensation and convection

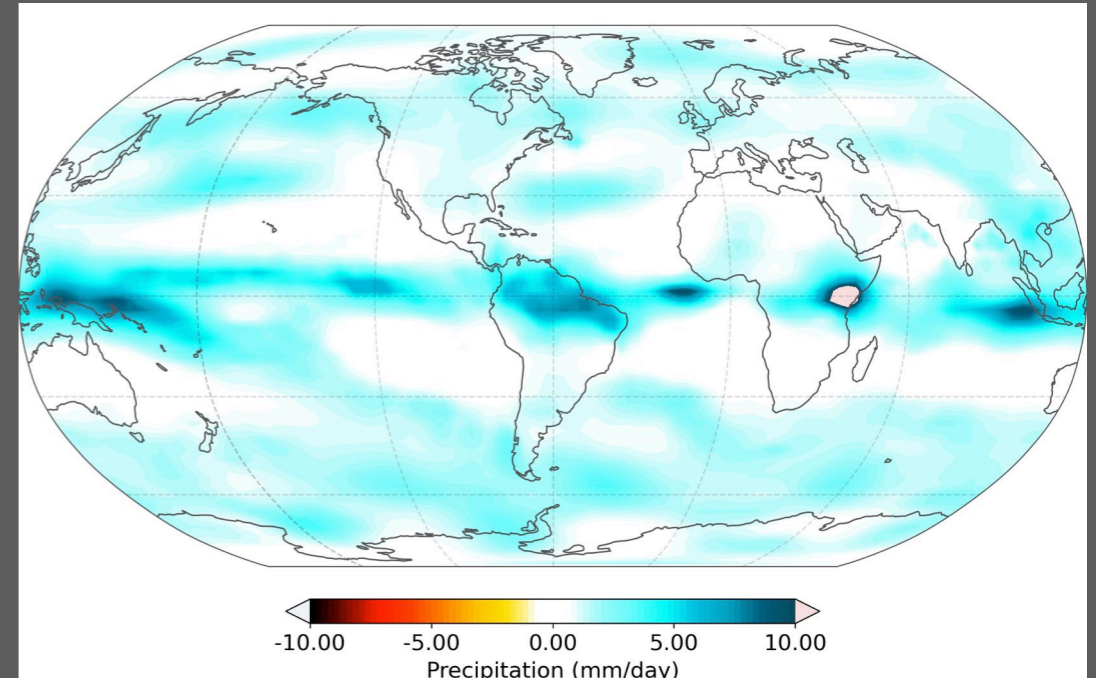
*DREAM can be run as a dry GCM with moisture acting as a passive tracer.*

*DREAM can also be run with active schemes for large scale rain and deep convection.*

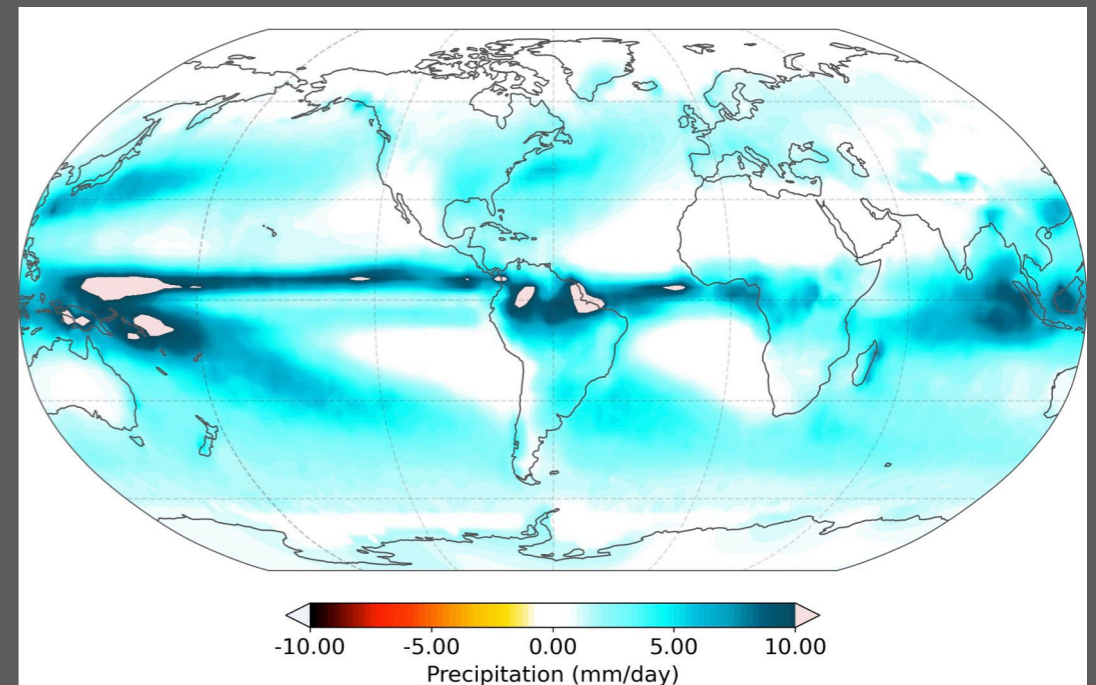
*The deep convection interacts with the dynamics through condensation heating in the upper troposphere.*

*Prescribed SST anomalies are converted into fixed anomalies in precipitation, which are also linked to anomalies in deep tropospheric heating.*

*This deep heating drives a dynamical response in the atmosphere, which in turn influences changes in precipitation over the continents.*



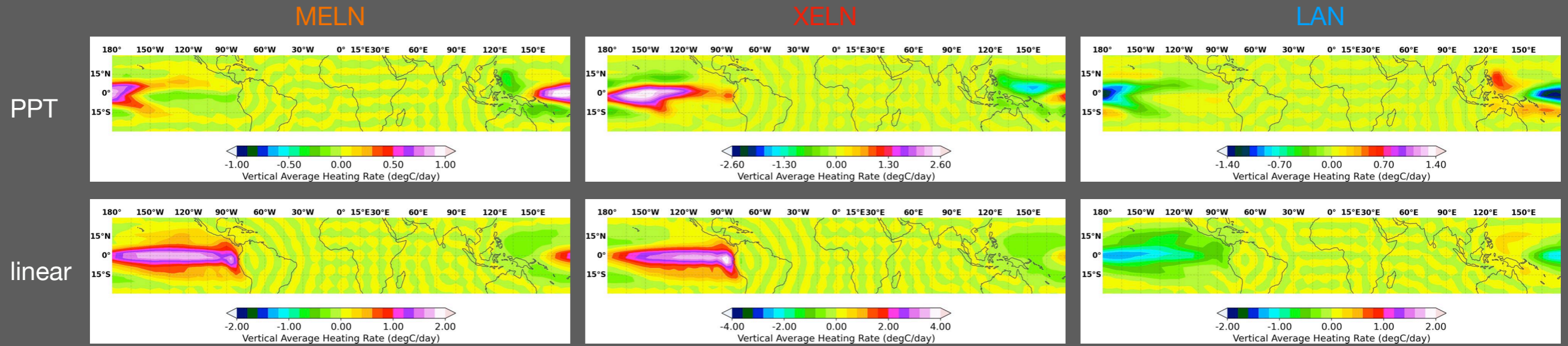
MAM Precipitation: DREAM



GPCP: 1991-2020

*Precipitation climatology from DREAM provides a reasonable basis for anomaly forecast experiments*

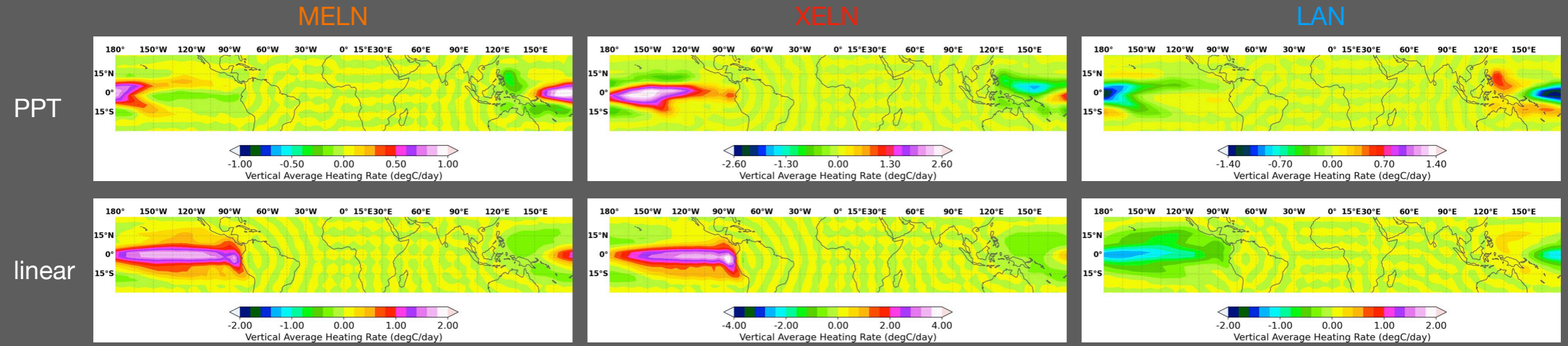
# From SSTA to precipitation



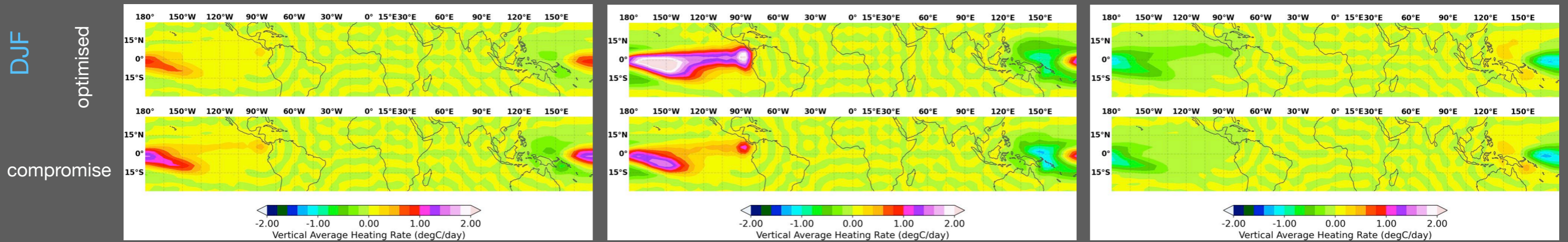
$$p' = M_{1FR} M_1 e^{\beta_{FR}\beta(T_C - T_0)} (D_C - D_0)(e^{\beta T_A} - 1)$$

	<b>RM1FR</b>	<b>beta</b>	<b>betaFR</b>	<b>Tref0</b>	<b>DIVref0</b>
tuned by eye	0.25	0.28	2.0	25	10
<b>MELN</b>	0.15	0.48	2.25	27	11
<b>XELN</b>	0.20	0.47	1.25	26	18
<b>LAN</b>	0.45	0.21	2.75	26	17
compromise	0.2	0.4	2.5	26.8	15

# From SSTA to precipitation



$$p' = M_{1FR} M_1 e^{\beta_{FR}\beta(T_C - T_0)} (D_C - D_0) (e^{\beta T_A} - 1)$$



Can we make this more intelligent ?